MICROWAVE LIMBSOUNDER HNO3 OBSERVATIONS: IMPLICATIONS FOR POLAR STRATOSPHERIC CLOUD COMPOSITIONS

Michelle L. Santee¹, Ross J. Salawitch¹, Gloria L. Manney¹, A zadeh Tabazadeh² William G. Read¹, Lucien Froidevaux¹, Joe W. Waters¹

Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA
NASA Ames Research Center, Moffett Field, CA

The catalytic destruction of O₃ in the polar lower stratosphere is initiated through heterogeneous chernical reactions on the surfaces of polar stratospheric cloud (PSC) particles. Accurate theoretical predictions of O₃ loss require knowledge of the phase and composition of these PSC particles. The canonical model of Type 1 PSCs [1] assumes, following thermodynamic stability arguments [2], that they are composed of solid nitric acid tril hydrate (NAT) particles. Recently, however, several observational, laboratory, and modeling studies [3] have suggested a liquid [1N():3/112S04/1 I₂O(ternary) solution as a more likely state for the Type I PSC particles.

To gain insight into the issue of PSC composition, we compare model results and measurements of gas-phase HNO3 made by the Microwaye Limb Sounder (M. LS) aboard the Upper Atmosphere Research Satellite (UARS) during the southern he misphere winters of 1992 and 1994. During the MLS south-looking period in late May/early June 1994, temperatures were below the threshold for the onset of NAT PSCs (~195 K) for two full weeks before MLS observed any reduction in gas-phase HNO₃ at 465 K (on 7 June, as temperatures approached the ice frost point). Loss of gas-phase H NO₃ continued throughout the MLS south-looking period; by the beginning of July there was essentially no gas-phase HNO3 remaining within the region of lowest, temperatures. A time series of MLS II NO3 mixing ratios averaged over the region of low temperatures on the 465 K surface is compared to predicted behavior assuming various models of 1'S(; formation. Specifically, we use NMC temperatures, an assumed water vapor mixing ratio of 4.5 PPmv and the formulae of Hanson and Mauersberger [2] and Worsnop et al. [4] to calculate the saturation vapor pressures of HNO3 over NAT and NAD (nitric acid dilty drate), respectively. Both sets of calculations predict low HNO3 vapor concentrations that do not correspond well with the M 1,S 1 INO3 observations. We also use an acrosol physical chemistry n nodel (APCM) [5] for the ternary system, assuming 10 ppbv of 1 INO₃ and background sulfate acrosol conditions. In this case, excellent agreement is found between the M LS measurements of gas-phase HNO3 and the HNO3 partitioning predicted by the APCM assuming a ternary solution composition for the PSCs.

These results can be contrasted to M 1,S IINO₃ observations from 1992 (although the observing period is not exactly the same, data exist for 2 June in 1992 and 1994). In both years the IINO₃ values on 2 June are highin the portion of the vortex where temperatures are not, low, due to descent of IINO₃-rich air from above. However, in 1992 there is a deficit in gas-phase IINO₃ coincident with the region where temperatures are below 195 K, conforming to the expected pattern for NAT PSC formation, whereas in 1994 IINO₃ values are high throughout the vortex, even inside the region of low temperatures. Trajectory calculations reveal a more rapid cooling rate for the parcels within the 195 1{ temperature contour in 1992 than in 1994 during the 20 days prior to 2 June. We are investigating the possible correlation between temperature history and PSC composition as inferred from M LS measurements of gas-phase IINO₃. We are also examining the differences between the 1992 and 1994 southern vortex M 1,S observations in relation to the changing sulfate aerosol loading as the stratosphere recovers from the Mt. Pinatubo eruption.

- [1] e.g., Turco et d., GRL 94, 16,493 (1989).
- [2] Hanson and Mauersberger, GRL 15, 855 (1988).
- [3] e.g., Zhang et al., J. Phys. Chem. 97, 8541 (1993); Carslaw et al., GRL 21, 2479 (1994); Drdla et al., GRL 21, 2475 (1994); Tabazadeh et al., GRL 21, 1619 (1994); Stefanutti et al., GRL 22, 2377 (1995); Toon and Tolbert, Nature 375, 218 (1995).
- [4] Worsnop et al., Science 259, 71 (1993).
- [5] Tabazadeh et (d., JGR 99, 12,897 (1994).